

Polarization characteristics of the atmosphere : Sensitivity study and variable information

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Background

- The polarization of the measured light depends on the optical system (mirror, grating, ...) and the actual incoming light (Schutgens and Stammes, 2003).
- Incorporating the polarization enables a more accurate retrieval of atmospheric properties and constituents (Mishchenko and Travis, 1997)
- Satellite groups measure the state of polarization to improve radiometric calibration.
 - **1.** *GOME* (*Burrows et al.*, 1999)
 - 2. SCIAMACHY (Bovensmann et al., 1999)
 - **3.** *GOME-2* (*Callies et al., 2000*)
- GEMS will develop a polarization correction algorithm using the RTM simulation results.

List of variables affecting the polarization

- Solar & Observation Geometry
- Surface reflectance
- Trace gases
- Aerosols
- Clouds
- etc.
- We plan to apply a real-time polarization correction using Look-Up-Table
- The table will be prepared based on the RTM simulation results

Method

Polarization Correction Algorithm(Sun and Xiong, 2007)

 $I' = hI\{1 + facos[2(\phi - \chi)]\}$

Polarization Term

I': Intensity reaching the CCD
h: Transmittance of the optical system
I: Intensity reaching the diffuser
a: Degree of polarization
φ: Angle of maximum transmission
χ: Phase angle of polarization w.r.t. instrument reference plane
f: Linear polarization sensitivity

Polarization : Trace Gases



Simulation is done for a molecular atmosphere using VLIDORT

- Residual of DoLP = DoLP (molec atm) DoLP(molec atm w/o X species)
- **DoLP** = $\frac{\sqrt{Q^2 + U^2}}{I}$
- Trace gas influence
 - $O_{3} >> O_{2} O_{2} > NO_{2}, SO_{2}, HCHO$
- Changes in the amount of ozone should be considered

Sensitivity : Ozone profile



- Ozone profile changes with latitude (especially, 10~20 km) and with season.
- DoLP depends on the ozone amount (and peak height).
- DoLP change is the largest around 308 nm.

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Sensitivity : Surface Reflectance

Molecular Atmosphere



• DoLP at 440 nm for various albedo (plot on SZA vs RAA)

- Relatively large DoLP occurs for SZA=90, RAA=90 when albedo is small

Sensitivity : Observation Geometry



Summer solstice for Seoul, 12 LST

DoLP at 440 nm as a function of SZA (plot on VZA vs RAA), with albedo=0.1
 DoLP tends to increase with VZA at small SZA, and large DoLP occurs for VZA=90 at large SZA.



DoLP at 440 nm for various RAA (plot on SZA vs VZA), with albedo=0.1
 When RAA is small, large DoLP occurs

Phase Function of Aerosols



Mixed aerosols (OPAC) and Dust aerosols



OPAC: CA(Continental Average), CC(Continental Clean), MP(Maritime Polluted) for RH=50%
 Dust: Asian Dust – knud, Saharan Dust – mitr(OPAC), dlike(dust-like)
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Polarization : Aerosol Type

AOD : 1.0 Albedo : 0.1





Polarization : AOD(Mineral Dust)



- DoLP changes with AOD
- DoLP changes for Adust and Saharan (MitR) dusts are different

Conclusion and Future Plan

- Polarization effects were simulated using RTM for GEMS polarization correction.
- Polarization is affected by observation geometry, trace gases, surface reflectivity and etc.
- The polarization due to the aerosol is different depending on the optical characteristics. On the other hands, it is less affected at shorter wavelengths.
- Additional analysis of aerosol and cloud polarization effects will be done.
- A polarization correction will be applied to improve accuracy of GEMS raw data.

Thank you ~

GEMS 측정 에너지 모의

| Time(hour) | Spring Equinox | | Summer Solstice | | Autumn Equinox | | Winter Solstice | |
|------------|------------------|-------------------|-----------------|----------------|----------------|----------------|-----------------|--------|
| | μ _o † | φ <mark>₀‡</mark> | μ | φ _o | μ | φ _o | μ | φ₀ |
| 06 | 91.18 | 88.83 | 76.20 | 70.82 | 88.16 | 91.06 | 103.68 | 109.21 |
| 07 | 79.46 | 98.05 | 64.76 | 78.89 | 76.60 | 100.40 | 92.68 | 117.69 |
| 08 | 67.92 | 108.00 | 52.99 | 87.20 | 65.19 | 110.70 | 82.65 | 127.12 |
| 09 | 57.05 | 119.68 | 41.12 | 96.76 | 54.61 | 123.01 | 73.95 | 137.95 |
| 10 | 47.53 | 134.39 | 29.57 | 109.80 | 45.62 | 138.73 | 67.01 | 150.54 |
| 11 | 40.49 | 153.55 | 19.33 | 132.57 | 39.45 | 159.06 | 62.50 | 164.91 |
| 12 | 37.43 | 177.02 | 14.13 | 178.45 | 37.60 | 183.07 | 60.98 | 180.43 |
| 13 | 39.30 | 201.11 | 18.85 | 225.62 | 40.67 | 206.44 | 62.67 | 195.92 |
| 14 | 45.50 | 221.47 | 28.95 | 249.30 | 47.71 | 225.54 | 67.33 | 210.20 |
| 15 | 54.52 | 237.18 | 40.48 | 262.64 | 57.21 | 240.23 | 74.39 | 222.70 |
| 16 | 65.12 | 249.48 | 52.34 | 272.32 | 68.07 | 251.90 | 83.17 | 233.44 |
| 17 | 76.54 | 259.77 | 64.12 | 280.66 | 79.61 | 261.85 | 93.29 | 242.79 |
| 18 | 88.10 | 269.11 | 75.58 | 288.73 | 91.35 | 271.07 | 104.30 | 251.25 |

† Solar Zenith Angle
‡ Solar Azimuth Angle

| Aerosol types | Components | N _i (cm ⁻³) | <i>M_i</i> (µg m⁻³) | Number mixing ratios (n) | Mass mixing ratios (m _j) |
|------------------------|--|--|--------------------------------------|--|--|
| Continental clean | total water soluble insoluble | 2600 2600 0.15 | 8.8 5.2 3.6 | 1.0 0.577E-4 | 0.591 0.409 |
| Continental average | total water soluble insoluble soot | 15 300 7000 0.4 8300 | 24.0 14.0 9.5 0.5 | 0.458 0.261E–4 0.542 | 0.583 0.396 0.021 |
| Continental polluted | total water soluble insoluble soot | 50 000 15 700 0.6 34 300 | 47.7 31.4 14.2 2.1 | 0.314 0.12E-4 0.686 | 0.658 0.298 0.044 |
| Urban | total water soluble insoluble soot | 158 000 28 000 1.5 130 000 | 99.4 56.0 35.6 7.8 | 0.177 0.949e–05 0.823 | 0.563 0.358 0.079 |
| Desert | total water soluble mineral (nuc.) mineral (acc.) mineral (coa.) | 2300 2000 269.5 30.5 0.142 | 225.8 4.0 7.5 168.7 45.6 | 0.87 0.117 0.133E–1 0.617–4 | 0.018 0.033 0.747 0.202 |
| Maritime clean | total water soluble sea salt (acc.) sea salt (coa.) | 1520 1500 20 3.2E–3 | 42.5 3.0 38.6 0.9 | 0.987 0.132E–1 0.211E–5 | 0.071 0.908 0.021 |
| Maritime polluted | total water soluble sea salt (acc.) sea salt (coa.) soot | 9000 3800 20 3.2E-3 5180 | 47.4 7.6 38.6 0.9 0.3 | 0.422 0.222E-2 0.356E-6 0.576 | 0.160 0.814 0.019 0.006 |